

GEOPHYSICS

Constraints on Earth's inner core composition inferred from measurements of the sound velocity of hcp-iron in extreme conditions

Tatsuya Sakamaki,^{1*} Eiji Ohtani,^{1,2} Hiroshi Fukui,^{3,4} Seiji Kamada,^{1,5} Suguru Takahashi,¹ Takanori Sakairi,¹ Akihiro Takahata,¹ Takeshi Sakai,⁶ Satoshi Tsutsui,⁷ Daisuke Ishikawa,⁷ Rei Shiraishi,¹ Yusuke Seto,⁸ Taku Tsuchiya,⁶ Alfred Q. R. Baron^{4,7}

Hexagonal close-packed iron (hcp-Fe) is a main component of Earth's inner core. The difference in density between hcp-Fe and the inner core in the Preliminary Reference Earth Model (PREM) shows a density deficit, which implies an existence of light elements in the core. Sound velocities then provide an important constraint on the amount and kind of light elements in the core. Although seismological observations provide density–sound velocity data of Earth's core, there are few measurements in controlled laboratory conditions for comparison. We report the compressional sound velocity (V_p) of hcp-Fe up to 163 GPa and 3000 K using inelastic x-ray scattering from a laser-heated sample in a diamond anvil cell. We propose a new high-temperature Birch's law for hcp-Fe, which gives us the V_p of pure hcp-Fe up to core conditions. We find that Earth's inner core has a 4 to 5% smaller density and a 4 to 10% smaller V_p than hcp-Fe. Our results demonstrate that components other than Fe in Earth's core are required to explain Earth's core density and velocity deficits compared to hcp-Fe. Assuming that the temperature effects on iron alloys are the same as those on hcp-Fe, we narrow down light elements in the inner core in terms of the velocity deficit. Hydrogen is a good candidate; thus, Earth's core may be a hidden hydrogen reservoir. Silicon and sulfur are also possible candidates and could show good agreement with PREM if we consider the presence of some melt in the inner core, anelasticity, and/or a premelting effect.

INTRODUCTION

Seismological data reveal important physical properties of Earth's core, such as density and elasticity. The core is mainly composed of iron (Fe), and the stable crystal structure of Fe under inner core conditions is hexagonal close-packed (hcp) (1). The density of the inner core in the Preliminary Reference Earth Model (PREM) (2), which is based on seismological data, is about 2 to 5% smaller than that of hcp-Fe (a core density deficit) (3), and it is accepted that the inner core consists of iron and light elements, such as silicon, sulfur, oxygen, hydrogen, and carbon (4).

The sound velocity (V_p) profile of Earth's core is important information obtained from seismological observations, and the density-velocity relation for Earth's core is well known. To constrain the abundances of light elements in Earth's core, Birch's law, a linear relation between sound velocity and density (5), has been used to extrapolate measured sound velocities to inner core conditions. To do this, the sound velocities of Fe and Fe alloys have been measured at high pressure and moderate temperatures (<2000 K) by inelastic x-ray scattering (IXS) (6–13) and nuclear inelastic scattering (NIS) (14–16). It is then desirable to extend these measurements as close as possible to core conditions. However, this is difficult, and sound velocity measurements

with static compression at temperatures above 2000 K have not been reported.

The sound velocity of pure hcp-Fe at high temperatures provides an important constraint on the composition of Earth's core: any model of the impurity composition of the core must both satisfy the density deficit from seismology and give an appropriate velocity. Here, we report the V_p of hcp-Fe up to 163 GPa and 3000 K (Fig. 1) based on a combination of the laser-heated diamond anvil cell (LHDAC) (17) and IXS measurements (18).

RESULTS

We fit our data by allowing the coefficients in the linear ρ - V_p relation to vary with temperature, which preserves Birch's law at any fixed temperature (Fig. 2A). A good fit to the data is possible by taking a first-order, linear, temperature dependence of the coefficients. The preferred orientation of the sample in the DAC is estimated using x-ray diffraction (XRD) pattern, and the velocity variation is less than 1.5% at 163 GPa and 3000 K (see also the discussion in the Supplementary Materials). We then parameterize the dependence as

$$V_p(\rho, T) = M\rho + B + A(T - T_0)(\rho - \rho^*) \quad (1)$$

We choose T_0 to be 300 K; M and B are the coefficients of Birch's law at room temperature, whereas A and ρ^* embody the temperature dependence. This form is convenient for the consideration of errors: the large data set at room temperature fixes M and B relatively precisely, whereas the sparse data at high temperatures are reflected in the uncertainty in A and ρ^* . The appearance of the crossover density ρ^* also explicitly indicates a maximum upper bound for the validity of this relation. Fitting our data, combined with shock compression data

¹Department of Earth and Planetary Materials Science, Graduate School of Science, Tohoku University, Sendai, Miyagi 980-8578, Japan. ²V.S. Sobolev Institute of Geology and Mineralogy, Siberian Branch, Russian Academy of Sciences, Novosibirsk, 630090, Russia. ³Center for Novel Material Science under Multi-Extreme Conditions, Graduate School of Material Science, University of Hyogo, Kamigori, Hyogo 678-1297, Japan. ⁴Materials Dynamics Laboratory, RIKEN SPring-8 Center, Sayo, Hyogo 679-5148, Japan. ⁵Frontier Research Institute for Interdisciplinary Sciences, Tohoku University, Sendai Miyagi 980-8578, Japan. ⁶Geodynamics Research Center, Ehime University, Matsuyama, Ehime 790-8577, Japan. ⁷Research and Utilization Division, Japan Synchrotron Radiation Research Institute, SPring-8, 1-1-1 Kouto, Sayo, Hyogo 679-5198, Japan. ⁸Department of Earth and Planetary Sciences, Kobe University, Kobe 657-8501, Japan.

*Corresponding author. E-mail: sakamaki@tohoku.ac.jp

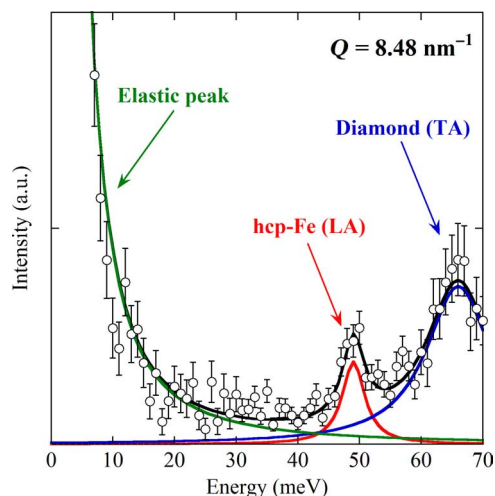


Fig. 1. IXS spectrum of hcp-Fe at 163 GPa and 3000 K. The peak at zero energy is from elastic scattering. Curves are individual contributions [green, elastic scattering; red, longitudinal acoustic (LA) phonons of hcp-Fe; blue, transverse acoustic (TA) phonons of diamond; fitting the experimental data with Lorentzian functions]. a.u., arbitrary units.

(19) along the Hugoniot, gives $M = 1.160 \pm 0.025$, $B = -3.43 \pm 0.29$, $A = 7.2 \times 10^{-5} \pm 3.6 \times 10^{-5}$, and $\rho^* = 14.2 \pm 1.5$. Details of the fitting are presented in the Supplementary Materials.

Figure 2B compares the present results to previous IXS measurements (6, 10) and ab initio calculations (20, 21). The present isothermal ρ - V_p fitting lines are consistent with previous IXS (6, 10). In the case of ab initio calculation, the lines have a good agreement with the results (20, 21) above 4000 K, although some discrepancies exist in the lower temperature range (<4000 K).

DISCUSSION

To demonstrate the impact of this work for understanding core composition, we focus on ρ and V_p at the inner core boundary (ICB), where the pressure has been determined to be 330 GPa by a seismological study (2). The recent discussion of the thermal state of the outer core based on the melting experiments of iron–light element systems suggests that the temperature at ICB is around 5500 K (22). The density of hcp-Fe can be calculated using the equation of state (3). The density-velocity relation from the present work extrapolated to this temperature is shown in Fig. 3. Also shown are the results of the PREM (2). The phenomenological Eq. 1 may be used for extrapolation to the inner core conditions (see also the discussion in the Supplementary Materials). As is apparent from Fig. 3, both the density and V_p of hcp-Fe are higher than PREM values: Earth's inner core has a 4 to 5% smaller density and a 4 to 10% smaller V_p than hcp-Fe at 5500 K. Thus, on the basis of the present sound velocity results, we conclude that light elements, or a combination of light elements and nickel, in the inner core decreases not only density but also simultaneously the V_p of hcp-Fe under inner core conditions. Assuming the lower core temperature model ($T_{\text{ICB}} = 4200$ K) proposed on the basis of the melting temperature of pyrolite (23), Earth's inner core has a 5 to 6% smaller density and a 7 to 13% smaller V_p than hcp-Fe at 4200 K.

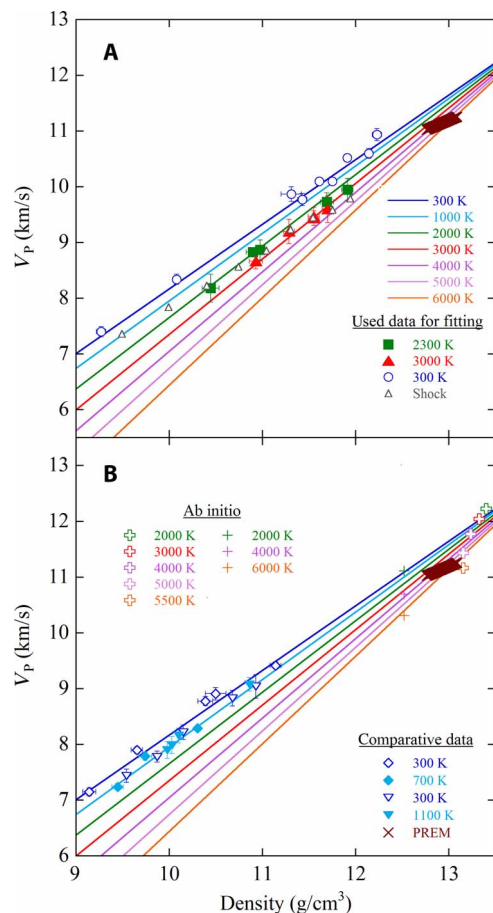


Fig. 2. Compressional velocity of hcp-Fe at several temperatures as a function of density. (A) Results of the present study and of Ohtani *et al.* (12) based on IXS and shock compression experiments as reported by Brown and McQueen (19). Isothermal ρ - V_p fitting lines are expressed as $V_p = [1.160\rho - 3.43] + [7.2 \times 10^{-5} \times (T - 300) \times (\rho - 14.2)]$. (B) Comparison of obtained isothermal ρ - V_p fitting lines with previous studies of NIS reported by Mao *et al.* (10), diamonds; IXS reported by Antonangeli *et al.* (6), inverted triangles; ab initio calculation reported by Vočadlo *et al.* (20), open crosses; and report of Sha and Cohen (21), crosses.

Several sound velocity and density measurements of pure iron and iron alloys demonstrate the effect of light elements on the properties of iron. The ρ - V_p plots of hcp-Fe (3, 12), hcp-Fe₉₂Ni₈ (14, 24), dhcp-Fe (11, 25), hcp-Fe₈₅Si₁₅ (10, 26), Fe₃S (13, 27), FeO (B1/rhombohedral phase) (7, 28), Fe₇C (9, 29), and Fe₇C₃ (30) are summarized in Fig. 4. The ρ - V_p plots of hcp-Fe at 4200 and 5500 K are also plotted in this diagram, and a star symbol indicates the ICB condition. Each ρ - V_p plot of iron alloys and iron–light element compounds is shown from the center of Earth (COE) to the ICB. The stars indicate the extrapolation to the ICB conditions ($T_{\text{ICB}} = 5500$ K) assuming that the temperature effect on these compounds is the same as that on hcp-Fe. This assumption is made on the basis of the few studies that show the temperature effects on iron alloys based on Birch's law. To account for the composition of Earth's inner core, the triangles (hcp-Fe–hcp-Fe₉₂Ni₈–Fe–light element compounds) need to overlap with PREM. In this discussion, we assumed Fe alloy to be an ideal mixture among various Fe-rich compounds, such as hcp-Fe, hcp-Fe–Ni, and Fe–light element

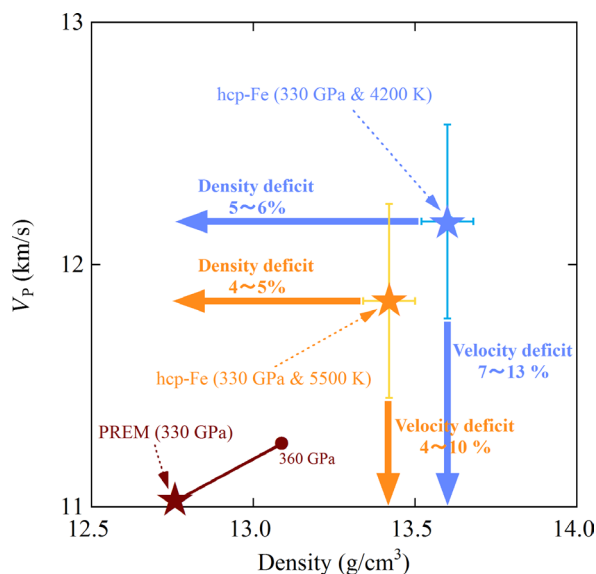


Fig. 3. Comparison of the ρ - V_p plot between hcp-Fe under ICB conditions and PREM. Stars indicate V_p and ρ of hcp-Fe and PREM at 330 GPa. The difference in the ρ - V_p plot between PREM and hcp-Fe at 5500 K (orange) shows a 4 to 5% core density deficit and a 4 to 10% core velocity deficit and that of 4200 K (blue) shows a 5 to 6% core density deficit and a 7 to 13% core velocity deficit.

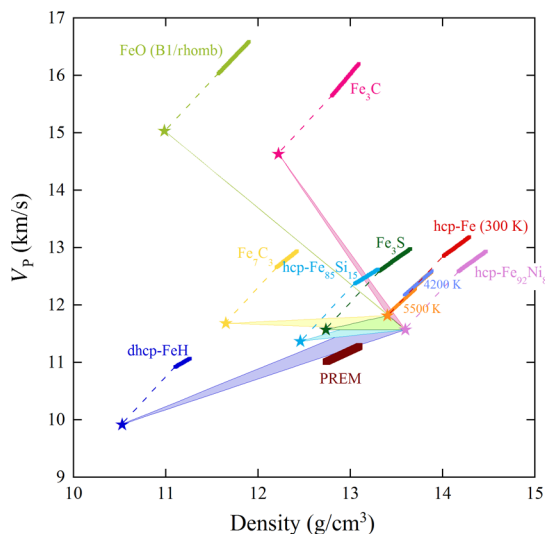


Fig. 4. Comparison of the ρ - V_p plot between hcp-Fe and Fe-light element compounds. The values of hcp-Fe and each Fe-light element compound [hcp-Fe₉₂Ni₈ (14, 24), purple; dhcp-FeH (11, 25), blue; hcp-Fe₈₅Si₁₅ (10, 26), aqua; Fe₇S (13, 27), green; FeO (B1/rhombohedral phase) (7, 28), yellow-green; Fe₃C (9, 29), pink; Fe₇C₃ (30), yellow] show distributions from the ICB to the COE. The obtained ρ - V_p of hcp-Fe at 5500 K is indicated by orange, and the star is at the expected ICB conditions. The other stars denote ρ - V_p of Fe-light element compounds under ICB conditions assuming that the temperature effects of these compounds are the same as those of hcp-Fe (the dashed lines are the temperature effect). Triangles connecting three stars (hcp-Fe-hcp-Fe₉₂Ni₈-Fe-light element compound) indicate the potential ρ - V_p region given by mixing three components.

alloys to estimate the average ρ - V_p of Fe alloy with light elements. Oxygen and carbon may not be suitable as major light elements in the core. On the other hand, silicon, sulfur, and hydrogen decrease both density and V_p and are thus potential candidates. Hydrogen is a good candidate because the mixing of hcp-Fe, hcp-Fe₉₂Ni₈, and dhcp-FeH can be accountable for the ρ - V_p plot of PREM. Thus, the inner core may be a hidden hydrogen reservoir in Earth. Silicon and sulfur can also meet an important requirement if their temperature effects on V_p are larger than that of hcp-Fe. In addition, the presence of melt in the inner core (31, 32), anelasticity (32), and/or a premelting effect (33), which causes the decrease in sound velocities, can support a silicon- or sulfur-rich inner core.

MATERIALS AND METHODS

Samples and cells

The IXS experiments and XRD experiments were conducted up to 163 GPa and 3000 K at the same beamline, BL35XU. High-pressure and high-temperature conditions were generated using a double-sided LHDAC with 150- and 250- μ m culets. The used laser heating system is COMPAT (17), which has been developed for an LHDAC in a limited experimental space, such as the beamline of a synchrotron radiation facility. Rhenium gaskets were preindented to 35 to 70 μ m in thickness, and a hole 50 to 100 μ m in diameter was drilled in the gasket as the sample chamber. The iron sample (99.99% purity) was compacted from a powder into a foil 45 to 90 μ m in diameter and 10 to 30 μ m in thickness. The sample was sandwiched between two NaCl layers that are 10 μ m in thickness, which were used as a pressure medium and a thermal insulator.

Laser heating

The average temperature of the heating spot obtained at 118 GPa and 3000 K is shown in fig. S1. The typical temperature uncertainty is 200 K due to fluctuation. The experimental duration at one pressure-temperature condition was about 4 to 6 hours. The temperature was monitored and controlled for 4 to 6 hours and recorded every 15 min during the heating experiments. Sample temperature under laser heating was determined by fitting Planck's formula to a spectrum of thermal radiation from a sample between 600 and 800 nm. The optics used to collect thermal radiation was calibrated with an intensity standard lamp. The hot spot is greater than the x-ray beam size at full width at half maximum (FWHM). We calculated the sample temperature by averaging the variation in the heating area probed by x-rays during the IXS measurements that lasted 4 to 6 hours.

IXS and XRD

The IXS and XRD experiments were performed at BL35XU of SPring-8 in Japan (18). We used the Si (9 9 9) backscattering optics, which provides an incident photon energy of 17.794 keV with an energy resolution of 2.8 meV at FWHM. The beam size was focused to 16 \times 16 μ m by a Kirkpatrick-Baez mirror pair for this work (34). The scattered x-rays were analyzed by 12 crystals, which are arranged in a two-dimensional (3 \times 4) array. The momentum transfer $Q = 2k_0 \sin(2\theta/2)$, where k_0 is the wave vector of the incident photons and 2θ is the scattering angle, was selected by rotating the spectrometer arm in the horizontal plane. Here, we collected IXS spectra in the range of $Q = 6.0$ to 10.0 nm^{-1} at each experimental pressure. The momentum resolution was set to about 0.4 nm^{-1} full width. IXS spectra of hcp-Fe at 163 GPa and 3000 K are shown in fig. S2.

The XRD patterns were obtained under the same experimental conditions as IXS using a flat panel area detector installed at BL35XU. An example of XRD pattern is shown in fig. S3. Rediffraction peaks may come from tails of x-ray beams. Pressure was calculated on the basis of XRD patterns using the equation of state of hcp-Fe (3). The lattice preferred orientation of this compressed sample was estimated using this diffraction pattern and discussed below. The experimental conditions are summarized in table S1.

Fits to the dispersion

Figure 1 shows the IXS spectra collected at 163 GPa and 3000 K. The spectra are characterized by an elastic contribution centered at zero energy and inelastic contributions from hcp-iron and diamond. The peak of the TA phonons of diamond appears at higher energies compared to hcp-iron because of the much higher sound velocity of diamond. The zero energy was well determined by the elastic peak. The energy positions of photons were extracted by fitting the data with a set of Lorentzian functions. The compressional sound velocity (V_p) was determined by fitting the phonon dispersion with a sine function

$$E[\text{meV}] = 4.192 \times 10^{-4} V_p[\text{m/s}] \times Q_{\text{MAX}}[\text{nm}^{-1}] \sin\left(\frac{\pi}{2} \frac{Q[\text{nm}^{-1}]}{Q_{\text{MAX}}[\text{nm}^{-1}]}\right) \quad (\text{S1})$$

where E and Q are the energy and the momentum of the acoustic mode, and Q_{MAX} is the first Brillouin zone edge (35). The expression can be derived from the solution of the dynamical matrix limited to the nearest-neighbor interaction within the framework of the Born-von Karman lattice dynamics theory (36). Here, Q_{MAX} was taken to be a free parameter in the fit, as in previous studies (9). V_p can change when Q_{MAX} is fixed to be $Qa/2\pi = 0.6$, but it remains within the error range for the fits to V_p when Q_{MAX} is free (fig. S4). Because our sample was polycrystalline, the determined velocity is an aggregate sound velocity, averaging over crystal orientations. We made critical assessments of texture effects and deviation from ideal average velocity as shown below. Figure S5 shows the dispersion curve at each experimental condition and the fits using eq. S1. The phonon energy of inelastic scattering increases with pressure, which indicates an increase of the compressional sound velocity with pressure. The results are summarized in table S2. Our 12 independent data points (albeit at mainly four different momentum transfers) provide a good constraint on the sound velocity. In addition, we conducted extra experiments to check the Q dependence of the obtained sound velocity. The difference in velocity at higher (IXS101) and lower (IXS102) Q in the same conditions is $\sim 0.35\%$, as shown in fig. S6.

Birch's law

Birch's law is a linear relation of the compressional sound velocity V_p with density ρ as $V_p = a\rho + b$ (4, 5), which arises in the physics of solids as a linearization of a power law (37). No regard is given to the temperature effect on the equation. To cover ρ - V_p data under a wide range of temperature conditions, the temperature effect on Birch's law should be considered. A recent ab initio calculation study (33) reports the premelting effect on sound velocities of hcp-Fe. Their results indicate that the temperature dependence of the velocity is

linear below melting temperature ($T/T_m < 0.9$). Assuming the linear temperature effect on the slope, a , and intercept, b , we modify Birch's law as

$$V_p[\text{km/s}] = \{a_0 + a_1 \cdot (T[\text{K}] - 300)\} \rho[\text{g/cm}^3] + \{b_0 + b_1 \cdot (T[\text{K}] - 300)\} \quad (\text{S2})$$

Equation S2 can be rewritten in the following form

$$V_p = [a_0\rho + b_0] + [a_1 \cdot (T - 300) (\rho + b_1/a_1)] \quad (\text{S3})$$

Here, we replaced a_0 with M , b_0 with B , and a_1 with A , and we refined $-b_1/a_1$ as ρ^*

$$V_p = [M\rho + B] + [A(T - 300) (\rho - \rho^*)] \quad (\text{S4})$$

Equation S4 can be divided into a temperature-independent term, $[M\rho + B_0]$, and a temperature-dependent term, $[A(T - 300)(\rho - \rho^*)]$. The first term shows typical Birch's law at 300 K. The second one exhibits the temperature effect on V_p at each density (pressure). At the same density, V_p decreases with temperature. In addition, the decrease in V_p at the same temperature becomes smaller with an increase in density. However, the linear temperature dependence implies that these relations should only be applied below a density, given by ρ^* in eq. S4: at ρ^* , V_p becomes independent of temperature, and above ρ^* , the temperature dependence becomes unphysical, with velocity increasing as temperature increases. ρ^* corresponds to a pressure of about 500 GPa, far beyond the center pressure of Earth.

To approximate our results and extrapolate them to different conditions, we used the linear form discussed above and simultaneously fit the IXS results at high temperatures reported here and at room temperature (12) and the shockwave data (19). The error bars of the shockwave data are not given precisely; hence, we consider a velocity error of 1 or 2%. Table S3 gives the resulting parameters. To check the validation of fitting parameters, we compared fit results with experimental data. Figure S7 shows that the fit results are within error bars of each experimental data; thus, the fitting parameters are reasonable. Notably, as is often the case for linear fits, the normalized covariance matrix of the fit parameters tends to be highly correlated, with diagonal values that approach unity magnitude. To estimate the error in the value at the core-mantle boundary (CMB), we therefore used the error propagation equation (38) with the full covariance matrix of the fit parameters, which, in this case, is

$$\begin{aligned} \sigma_{V_p, V_p} = & \sigma_{MM} \left(\frac{\partial V_p}{\partial M} \right)^2 + \sigma_{BB} \left(\frac{\partial V_p}{\partial B} \right)^2 + \sigma_{AA} \left(\frac{\partial V_p}{\partial A} \right)^2 + \\ & \sigma_{\rho^* \rho^*} \left(\frac{\partial V_p}{\partial \rho^*} \right)^2 + 2\sigma_{MB} \left(\frac{\partial V_p}{\partial M} \right) \left(\frac{\partial V_p}{\partial B} \right) + \\ & 2\sigma_{MA} \left(\frac{\partial V_p}{\partial M} \right) \left(\frac{\partial V_p}{\partial A} \right) + 2\sigma_{M\rho^*} \left(\frac{\partial V_p}{\partial M} \right) \left(\frac{\partial V_p}{\partial \rho^*} \right) + \\ & 2\sigma_{AB} \left(\frac{\partial V_p}{\partial A} \right) \left(\frac{\partial V_p}{\partial B} \right) + 2\sigma_{B\rho^*} \left(\frac{\partial V_p}{\partial B} \right) \left(\frac{\partial V_p}{\partial \rho^*} \right) + \\ & 2\sigma_{A\rho^*} \left(\frac{\partial V_p}{\partial A} \right) \left(\frac{\partial V_p}{\partial \rho^*} \right) \end{aligned} \quad (\text{S5})$$

where σ_{AB} is the covariance of A and B and the error in the velocity $V(\rho, T)$ is given by $\sigma = \sqrt{\sigma_{V_P V_P} |_{\rho, T}}$. The results given in table S3 suggest that the velocity determination is fairly robust to small changes in the errors of the shockwave data.

Fitting using our high-temperature data (2300 and 3000 K) and shock compression data provided the parameters of eq. S4 in table S3. However, in the case of the combination of ambient temperature data (300 K) and shock compression data, fitting was not completed successfully. The fitting curves in Fig. 2 were obtained from the data at 2300, 3000 (this study), and 300 K (12) and from shock data (19).

We checked the fitting by a power law as suggested by Mao *et al.* (10). Our analyses revealed that the fitting by the power law shows almost no curvature and is very close to the linear extrapolation, suggesting that our linear extrapolation is valid.

Equation 1 gives the isothermal ρ - V_P relation by plugging the temperature. To compare with PREM, we assumed that the CMB temperature is 5500 K (22). The density at 330 GPa and 5500 K was calculated to be $13.42 \pm 0.08 \text{ g/cm}^3$ using the equation of state of hcp-Fe (2). Thus, the velocity of hcp-Fe at 330 GPa and 5500 K is $11.85 \pm 0.40 \text{ km/s}$. In the case of 5000 K, $\rho_{330\text{GPa}, 5000\text{K}} = 13.49 \pm 0.08 \text{ g/cm}^3$ and $V_{P, 330\text{GPa}, 5000\text{K}} = 11.98 \pm 0.40 \text{ km/s}$, whereas at 6000 K, $\rho_{330\text{GPa}, 6000\text{K}} = 13.34 \pm 0.08 \text{ g/cm}^3$, that is, $V_{P, 330\text{GPa}, 6000\text{K}} = 11.69 \pm 0.40 \text{ km/s}$. The ρ and V_P of PREM are 12.764 cm^3 and 11.028 km/s , respectively (2). In the temperature region between 5000 and 6000 K, the hcp-Fe indicates larger density and compressional sound velocity than PREM at 330 GPa. Therefore, the density and velocity deficits in Earth's inner core compared to hcp-Fe are $[\Delta\rho_{(\text{hcp-Fe} - \text{PREM})}/\rho_{\text{hcp-Fe}}]_{5500\text{K}} = 4$ to 5% and $[\Delta V_{P(\text{hcp-Fe} - \text{PREM})}/V_{P(\text{hcp-Fe})}]_{5500\text{K}} = 4$ to 10%, respectively.

In an analysis for the conclusion, we assumed that the temperature effect on Birch's law for all iron-light element compounds is the same as that for hcp-Fe. The present conclusion on the light elements in the inner core does not change without unusual reduction of the sound velocity of FeO and Fe₃C with increasing temperature, which is not likely (39–41).

Ab initio calculation

To evaluate the effect of the preferred orientation on the sound velocity measured in this work, we calculated the anisotropy of the elastic constants of hcp-Fe under the conditions of the present work. We calculated the elastic constants of hcp-Fe at 163 GPa and 3000 K based on an ab initio calculation: 1321.5 GPa for C_{11} , 661.3 GPa for C_{12} , 620.4 GPa for C_{13} , 1419.7 GPa for C_{33} , and 211.7 GPa for C_{44} . The C_{ij} calculated here at 163 GPa and 3000 K is consistent with previous work under similar conditions. Details of the ab initio calculation are given by Tsuchiya and Fujibuchi (42). The difference between C_{11} and C_{33} of hcp-Fe becomes smaller with temperature because C_{11} is 1552.8 GPa and C_{33} is 1767.0 GPa at 180 GPa and 0 K. The difference will become much smaller at higher temperatures and hcp-Fe will become isotropic, as reported by Sha and Cohen (21).

Preferred orientation

We need to consider the effect of the tendency of polycrystalline iron to preferentially align with the c axis parallel to the compressional axis under nonhydrostatic conditions (43). The preferred orientation was calculated using the XRD pattern at 163 GPa and 3000 K given in fig. S3. The preferred orientation of the sample is shown in fig. S8, which indicates that the compressed sample has a concentration of the c axis to the compression direction. (The concentration scale is 10, whereas the scale for 100% concentration of the axis to the compression di-

rection is 2000.) The preferred orientation analysis of the powder diffraction pattern was made on the basis of the procedure by Seto (44) and Niwa *et al.* (45). A pole figure (density map of a crystallographic axis) can be calculated from the diffraction pattern. The sound velocity distribution of the polycrystalline aggregate shown in fig. S8 was calculated using C_{ij} given above based on the program Anis2k (46). The midrange of this velocity distribution is almost the same as the Voigt-Reuss-Hill average. The resultant anisotropy of P -wave velocity is given in fig. S9, indicating that the sound velocity anisotropy calculated for the preferred orientation in the present DAC experiment at 163 GPa and 3000 K is less than 1.3%. Therefore, we do not expect to have a significant impact on this result. The anisotropy of the sound velocity was estimated by using C_{ij} calculated for hcp-Fe using the procedure by Tsuchiya and Fujibuchi (42).

SUPPLEMENTARY MATERIALS

Supplementary material for this article is available at <http://advances.sciencemag.org/cgi/content/full/2/2/e1500802/DC1>

- Fig. S1. Typical temperature profiles across the laser-heated spot obtained in IXS69.
 Fig. S2. IXS spectra of hcp-Fe at 163 GPa and 3000 K.
 Fig. S3. XRD pattern at 163 GPa and 3000 K.
 Fig. S4. Comparison of dispersion curves determined by fitting with Q_{MAX} free and Q_{MAX} fixed.
 Fig. S5. Longitudinal acoustic phonon dispersion of hcp-Fe at different experimental conditions.
 Fig. S6. Comparison of dispersion curve between IXS101 and IXS102.
 Fig. S7. Comparison of experimental data with fit result.
 Fig. S8. Pole figures for hcp-Fe at 163 GPa and 3000 K.
 Fig. S9. P -wave velocity anisotropy of hcp-Fe at 163 GPa and 3000 K based on the diffraction pattern.
 Table S1. IXS experimental conditions.
 Table S2. Results of density and compressional sound velocity measurements.
 Table S3. Results of different fits to the data sets as described in the text.

REFERENCES AND NOTES

- S. Tateno, K. Hirose, Y. Ohishi, Y. Tatsumi, The structure of iron in Earth's inner core. *Science* **330**, 359–361 (2010).
- A. M. Dziewonski, D. L. Anderson, Preliminary reference Earth model. *Phys. Earth Planet. Inter.* **25**, 297–356 (1981).
- A. Dewaele, P. Loubeyre, F. Occelli, M. Mezouar, P. I. Dorogokupets, M. Torrent, Quasihydrostatic equation of state of iron above 2 Mbar. *Phys. Rev. Lett.* **97**, 215504 (2006).
- F. Birch, Elasticity and constitution of the Earth's interior. *J. Geophys. Res.* **57**, 227–286 (1952).
- F. Birch, Composition of the Earth's mantle. *Geophys. J. Int.* **4**, 295–311 (1961).
- D. Antonangeli, T. Komabayashi, F. Occelli, E. Borissenko, A. C. Walters, G. Fiquet, Y. Fei, Simultaneous sound velocity and density measurements of hcp iron up to 93 GPa and 1100 K: An experimental test of the Birch's law at high temperature. *Earth Planet. Sci. Lett.* **331–332**, 210–214 (2012).
- J. Badro, G. Fiquet, F. Guyot, E. Gregoryanz, F. Occelli, D. Antonangeli, M. d'Astuto, Effect of light elements on the sound velocities in solid iron: Implications for the composition of Earth's core. *Earth Planet. Sci. Lett.* **254**, 233–238 (2007).
- G. Fiquet, J. Badro, F. Guyot, H. Requardt, M. Krisch, Sound velocities in iron to 110 gigapascals. *Science* **291**, 468–471 (2001).
- G. Fiquet, J. Badro, E. Gregoryanz, Y. Fei, F. Occelli, Sound velocity in iron carbide (Fe₃C) at high pressure: Implications for the carbon content of the Earth's inner core. *Phys. Earth Planet. Inter.* **172**, 125–129 (2009).
- Z. Mao, J.-F. Lin, J. Liu, A. Alatas, L. Gao, J. Zhao, H.-K. Mao, Sound velocities of Fe and Fe-Si alloy in the Earth's core. *Proc. Natl. Acad. Sci. U.S.A.* **109**, 10239–10244 (2012).
- Y. Shibazaki, E. Ohtani, H. Fukui, T. Sakai, S. Kamada, D. Ishikawa, S. Tsutsui, A. Q. R. Baron, N. Nishitani, N. Hirao, K. Takemura, Sound velocity measurements in dhcp-FeH up to 70 GPa with inelastic X-ray scattering: Implications for the composition of the Earth's core. *Earth Planet. Sci. Lett.* **313–314**, 79–85 (2012).
- E. Ohtani, Y. Shibazaki, T. Sakai, K. Mibe, H. Fukui, S. Kamada, T. Sakamaki, Y. Seto, S. Tsutsui, A. Q. R. Baron, Sound velocity of hexagonal close-packed iron up to core pressures. *Geophys. Res. Lett.* **40**, 5089–5094 (2013).

13. S. Kamada, E. Ohtani, H. Fukui, T. Sakai, H. Terasaki, S. Takahashi, Y. Shibazaki, S. Tsutsui, A. Q. R. Baron, N. Hirao, Y. Ohishi, The sound velocity measurements of Fe₃S. *Am. Mineral.* **99**, 98–101 (2014).
14. J.-F. Lin, V. V. Struzhkin, W. Sturhahn, E. Huang, J. Zhao, M. Y. Hu, E. E. Alp, H.-K. Mao, N. Boctor, R. J. Hemley, Sound velocities of iron-nickel and iron-silicon alloys at high pressures. *Geophys. Res. Lett.* **30**, 2112 (2003).
15. J.-F. Lin, W. Sturhahn, J. Zhao, G. Shen, H.-k. Mao, R. J. Hemley, Sound velocities of hot dense iron: Birch's law revisited. *Science* **308**, 1892–1894 (2005).
16. H. K. Mao, J. Xu, V. V. Struzhkin, J. Shu, R. J. Hemley, W. Sturhahn, M. Y. Hu, E. E. Alp, L. Vocadlo, D. Alfè, G. D. Price, M. J. Gillan, M. Schwoerer-Böhning, D. Häusermann, P. Eng, G. Shen, H. Giefers, R. Lübbers, G. Wortmann, Phonon density of states of iron up to 153 gigapascals. *Science* **292**, 914–916 (2001).
17. H. Fukui, T. Sakai, T. Sakamaki, S. Kamada, S. Takahashi, E. Ohtani, A. Q. R. Baron, A compact system for generating extreme pressures and temperatures: An application of laser-heated diamond anvil cell to inelastic X-ray scattering. *Rev. Sci. Instrum.* **84**, 113902 (2013).
18. A. Q. R. Baron, Y. Tanaka, S. Goto, K. Takeshita, T. Matsushita, T. Ishikawa, An X-ray scattering beamline for studying dynamics. *J. Phys. Chem. Solids* **61**, 461–465 (2000).
19. J. M. Brown, R. G. McQueen, Phase transitions, Grüneisen parameter, and elasticity for shocked iron between 77 GPa and 400 GPa. *J. Geophys. Res.* **91**, 7485–7494 (1986).
20. L. Vočadlo, D. P. Dobson, I. G. Wood, *Ab initio* calculations of the elasticity of hcp-Fe as a function of temperature at inner-core pressure. *Earth Planet. Sci. Lett.* **288**, 534–538 (2009).
21. X. Sha, R. E. Cohen, Elastic isotropy of ϵ -Fe under Earth's core conditions. *Geophys. Res. Lett.* **37**, L10302 (2010).
22. H. Terasaki, S. Kamada, T. Sakai, E. Ohtani, N. Hirao, Y. Ohishi, Liquidus and solidus temperatures of a Fe–O–S alloy up to the pressures of the outer core: Implication for the thermal structure of the Earth's core. *Earth Planet. Sci. Lett.* **304**, 559–564 (2011).
23. R. Nomura, K. Hirose, K. Uesugi, Y. Ohishi, A. Tsuchiyama, A. Miyake, Y. Ueno, Low core-mantle boundary temperature inferred from the solidus of pyrolite. *Science* **343**, 522–525 (2014).
24. T. Sakai, S. Takahashi, N. Nishitani, I. Mashino, E. Ohtani, N. Hirao, Equation of state of pure iron and Fe_{0.9}Ni_{0.1} alloy up to 3 Mbar. *Phys. Earth Planet. Inter.* **228**, 114–126 (2014).
25. N. Hirao, T. Kondo, E. Ohtani, K. Takemura, T. Kikegawa, Compression of iron hydride to 80 GPa and hydrogen in the Earth's inner core. *Geophys. Res. Lett.* **31**, L06616 (2004).
26. J.-F. Lin, A. J. Campbell, D. L. Heinz, G. Shen, Static compression of iron-silicon alloys: Implications for silicon in the Earth's core. *J. Geophys. Res.* **108**, 2045 (2003).
27. C. T. Seagle, A. J. Campbell, D. L. Heinz, G. Shen, V. B. Prakapenka, Thermal equation of state of Fe₃S and implications for sulfur in the Earth's core. *J. Geophys. Res.* **111**, B06209 (2006).
28. R. A. Fischer, A. J. Campbell, G. A. Shofner, O. T. Lord, P. Dera, V. B. Prakapenka, Equation of state and phase diagram of FeO. *Earth Planet. Sci. Lett.* **304**, 496–502 (2011).
29. N. Sata, K. Hirose, G. Shen, Y. Nakajima, Y. Ohishi, N. Hirao, Compression of FeSi, Fe₃C, Fe_{0.95}O, and FeS under the core pressures and implication for light element in the Earth's core. *J. Geophys. Res.* **115**, B09204 (2010).
30. C. Prescher, L. Dubrovinsky, E. Bykova, I. Kupenko, K. Glazyrin, A. Kantor, C. McCammon, M. Mookherjee, Y. Nakajima, N. Miyajima, R. Sinmyo, V. Cerantola, N. Dubrovinskaia, V. Prakapenka, R. Rüffer, A. Chumakov, M. Hanfland, High Poisson's ratio of Earth's inner core explained by carbon alloying. *Nat. Geosci.* **8**, 220–223 (2015).
31. S. C. Singh, M. A. J. Taylor, J. P. Montagner, On the presence of liquid in Earth's inner core. *Science* **287**, 2471–2474 (2000).
32. L. Vočadlo, *Ab initio* calculations of the elasticity of iron and iron alloys at inner core conditions: Evidence for a partially molten inner core? *Earth Planet. Sci. Lett.* **254**, 227–232 (2007).
33. B. Martorell, L. Vočadlo, J. Brodholt, I. G. Wood, Strong premelting effect in the elastic properties of hcp-Fe under inner-core conditions. *Science* **342**, 466–468 (2013).
34. D. Ishikawa, H. Uchiyama, S. Tsutsui, H. Fukui, A. Q. R. Baron, Compound focusing for hard-x-ray inelastic scattering. *Proc. SPIE* 8848–88480F (2013).
35. G. Fiquet, J. Badro, F. Guyot, Ch. Bellin, M. Krisch, D. Antonangeli, H. Requardt, A. Mermet, D. Farber, C. Aracne-Ruddle, J. Zhang, Application of inelastic X-ray scattering to the measurements of acoustic wave velocities in geophysical materials at very high pressure. *Phys. Earth Planet. Inter.* **143–144**, 5–18 (2004).
36. N. W. Ashcroft, N. D. Mermin, *Solid State Physics* (Saunders College Publishing & Harcourt Brace College Publishers, Orlando, FL, 1976) p. 826.
37. D. H. Chung, Birch's law: Why is it so good? *Science* **177**, 261–263 (1972).
38. P. R. Bevington, D. K. Robinson, *Data Reduction and Error Analysis for the Physical Sciences* (McGraw-Hill, New York, ed. 3, 2003), pp. 36–50.
39. R. Jeanloz, T. J. Ahrens, Equation of state of FeO and CaO. *Geophys. J. R. Astron. Soc.* **62**, 505–528 (1980).
40. T. Yagi, K. Fukuoka, H. Takei, Y. Syono, Shock compression of wüstite. *Geophys. Res. Lett.* **15**, 816–819 (1988).
41. L. Gao, B. Chen, J. Zhao, E. E. Alp, W. Sturhahn, J. Li, Effect of temperature on sound velocities of compressed Fe₃C, a candidate component of the Earth's inner core. *Earth Planet. Sci. Lett.* **309**, 213–220 (2011).
42. T. Tsuchiya, M. Fujibuchi, Effects of Si on the elastic property of Fe at Earth's inner core pressures: First principles study. *Phys. Earth Planet. Inter.* **174**, 212–219 (2009).
43. D. Antonangeli, F. Occelli, H. Requardt, J. Badro, G. Fiquet, M. Krisch, Elastic anisotropy in textured hcp-iron to 112 GPa from sound wave propagation measurements. *Earth Planet. Sci. Lett.* **225**, 243–251 (2004).
44. Y. Seto, Whole pattern fitting for two-dimensional diffraction patterns from polycrystalline materials. *Rev. High Pressure Sci. Technol.* **22**, 144–152 (2012).
45. K. Niwa, N. Miyajima, Y. Seto, K. Ohgushi, H. Gotou, T. Yagi, *In situ* observation of shear stress-induced perovskite to post-perovskite phase transition in CaIrO₃ and the development of its deformation texture in a diamond-anvil cell up to 30 GPa. *Phys. Earth Planet. Inter.* **194–195**, 10–17 (2012).
46. D. Mainprice, A FORTRAN program to calculate seismic anisotropy from the lattice preferred orientation of minerals. *Comput. Geosci.* **16**, 385–393 (1990).

Acknowledgments: This study was conducted as part of the Global Center of Excellence Program of Earth and Planetary Science, Tohoku University, Japan. The synchrotron radiation experiments were performed under contract of the SPring-8 (proposal nos. 2012A1255, 2012B1439, 2013A1377, 2013A1492, 2013B1078, 2013B1094, 2014A1100, 2014B1269, and 2014B1465). **Funding:** This work was supported by a Grant-in-Aid for Specially Promoted Research (no. 22000002) and Scientific Research (no.15H05748) to E.O. and Scientific Research to T. Sakamaki (no. 24840004 and no. 25800292) from the Ministry of Education, Culture, Sports, Science, and Technology of the Japanese Government. This work was also supported partly by the Ministry of Education and Science of the Russian Federation (project 14.B25.31.0032 to E.O.). **Author contributions:** T. Sakamaki, E.O., H.F., S.K., S. Takahashi, T. Sakairi, A.T., T. Sakai, S. Tsutsui, D.I., and A.Q.R.B. performed the IXS and XRD experiments. R.S., Y.S., and T.T. carried out several calculations. T. Sakamaki, E.O., H.F., and A.Q.R.B. made arguments on geological applications of the experimental results and prepared the paper. All authors discussed the results and commented on the manuscript. **Competing interests:** The authors declare that they have no competing interests. **Data and materials availability:** All data needed to evaluate the conclusions in the paper are present in the paper and/or the Supplementary Materials. Additional data related to this paper may be requested from the authors.

Submitted 18 June 2015
Accepted 18 December 2015
Published 26 February 2016
10.1126/sciadv.1500802

Citation: T. Sakamaki, E. Ohtani, H. Fukui, S. Kamada, S. Takahashi, T. Sakairi, A. Takahata, T. Sakai, S. Tsutsui, D. Ishikawa, R. Shiraiishi, Y. Seto, T. Tsuchiya, A. Q. R. Baron, Constraints on Earth's inner core composition inferred from measurements of the sound velocity of hcp-iron in extreme conditions. *Sci. Adv.* **2**, e1500802 (2016).

Constraints on Earth's inner core composition inferred from measurements of the sound velocity of hcp-iron in extreme conditions

Tatsuya Sakamaki, Eiji Ohtani, Hiroshi Fukui, Seiji Kamada, Suguru Takahashi, Takanori Sakairi, Akihiro Takahata, Takeshi Sakai, Satoshi Tsutsui, Daisuke Ishikawa, Rei Shiraishi, Yusuke Seto, Taku Tsuchiya and Alfred Q. R. Baron

Sci Adv 2 (2), e1500802.

DOI: 10.1126/sciadv.1500802

ARTICLE TOOLS

<http://advances.sciencemag.org/content/2/2/e1500802>

SUPPLEMENTARY MATERIALS

<http://advances.sciencemag.org/content/suppl/2016/02/23/2.2.e1500802.DC1>

REFERENCES

This article cites 43 articles, 10 of which you can access for free
<http://advances.sciencemag.org/content/2/2/e1500802#BIBL>

PERMISSIONS

<http://www.sciencemag.org/help/reprints-and-permissions>

Use of this article is subject to the [Terms of Service](#)

Science Advances (ISSN 2375-2548) is published by the American Association for the Advancement of Science, 1200 New York Avenue NW, Washington, DC 20005. The title *Science Advances* is a registered trademark of AAAS.

Copyright © 2016, The Authors